Martian Groundwater Outflows in Flume Experiments Processes and Morphological Properties

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INTRODUCTION

- There are many channels on Mars, but climate conditions were different than on Earth.
- Different sources of water have been proposed for Mars, including groundwater as main source for channel formation [1,2,3].

AIMS

- Knowledge on groundwaterinduced channels is minimal to limited occurence on Earth
- We aim to extend the knowled processes and resulting morph systems from scaled flume exp



	METHODS	prec
	• Experimental setup consists of a flume of	• Press
due	6 m long x 4 m wide and 1.20 m deep.	uran
٦.	• Simulation of seepage from sub-	– SU
dge on related	surface groundwater level from a distant	- SU
hology for these	source using a constant head tank.	lit
periments.	• Seepage from a local source (e.g. melt or	• Data





EXPERIMENT MOVIES

http://goo.gl/gfUbO

cipitation) was simulated by rain simulators. ssurized aquifer release using a subsurface inage pipe with forced discharge, at: ub-lithostatic pressure (only seepage) uper-lithostatic pressure (sediment ifted by water pressure) ta: time-lapse imagery and laserscan DEMs. MARS LOOK-ALIKE **KEY FEATURES** • Different sizes of valleys

haped heads

- due to flow piracy. Theater-shaped valley heads due to mass wasting processes.
- Valley depth relates to groundwater level.
- Further developed valleys are deeper as groundwater level is deeper upstream.

• Several valleys similar in size, due to absence of flow piracy.

- Headward development by mass wasting.
- Shallow valleys, due to high groundwater level.
- Simulated in experiment as precipitation, but could be melt of snow or subsurface ice.
- Converging flow features upstream: feathershaped head.
- Deposition of lobes after first overflow due to infiltration in unsaturated substrate (sieve deposits).
- No morphology left by actual seepage process.
- Not found on Mars without pits or chaos (see next).
- Similar features as sublithostatic pressure, but:
- Cracks and breaking of surface due to superlithostatic pressure.
- Pits in source area carved by emerging groundwater.
- Converging flow features disconnected from source area.



Fig. 1 Valley profiles

- as valleys develop (Fig. 4a). In the local case (Fig. 5a), the rate remains fairly constant. - Valley lengthening slows in both types of experiments (Fig. 4b, 5b). • Erosion takes place in pulses, which are more sudden in the distal cases (Fig. 4d) due to the collapsing nature of the headward development and widening.



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MORPHOLOGICAL ANALYSIS (SAPPING ONLY)

Local N

• Sapping valleys fed by distal groundwater source are deeper and have more pronounced valley heads (Fig. 1).

Valley profile

0.5

Initital surface

- In both cases, valleys are steeper in the upstream part (Fig. 2). This relates to the difference in processes: mudflows in the upstream end, fluvial transport downstream. • Valleys become more U-shaped when
- they develop (Fig. 3). Valleys fed by distal groundwater have a higher shape index, as the valleys have steeper cliffs.





Fig. 3 Valley shapes

MORPHOLOGICAL DEVELOPMENT (SAPPING ONLY)

• Valleys become wider, deeper and longer during the experiments. - In the distal cases, widening slows

1.5

Distance along valley (m)

• In the distal experiments, the number of active valleys decreased, due to groundwater piracy.



Fig. 4 Morphological development distal sapping experiments.



Fig. 5 Morphological development local sapping experiments.

CONCLUSIONS

- Different sources of groundwater for channel formation produce distinct types of valleys and channels.
- Groundwater sapping:
- Produces theater-shaped valley heads.
- Flow piracy occurs when the water source is distal, this focusses flow and enhances development of a few channels.
- Two processes, mudflow and fluvial flow are shown by a break in slope.
- Erosion takes place in pulses due to the collapsing development.
- Pressurized groundwater release:
- Results in channel head with converging flow features.
- Downstream lobate deposits on unsaturated sediment.
- Super-lithostatic pressure breaks surface and forms pits in the source area.

References [1] Howard A.D. & McLane C.F. (1988) WRR 24(1), 1659-1674. [2] Kite E.S. et al. (2011) JGR 116, E07002. [3] Andrews-Hanna J.C. & Phillips R.J. (2007) JGR 112, E08001. Image credits HiRISE: NASA/JPL/University of Arizona, THEMIS: NASA/JPL/ ASU. Funding WAM is supported by NWO grant ALW-GO-PL/10-01 to MGK.