

1 - INTRODUCTION

The occurrence of salt diapirs is strongly associated with potential geothermal and hydrocarbon energy sources. One occurrence of salt diapirism is commonly seen in areas undergoing thin skinned extension (e.g. North Angolan passive margin). Although this feature has been the subject of a few structural geological and analogue studies, research into the dynamics of such a system is scarce. Therefore we applied two-dimensional numerical modelling to address the following question:

Which geological/geometrical or material parameters affect the growth rate and shape of the diapir and how?

2 - METHODOLOGY

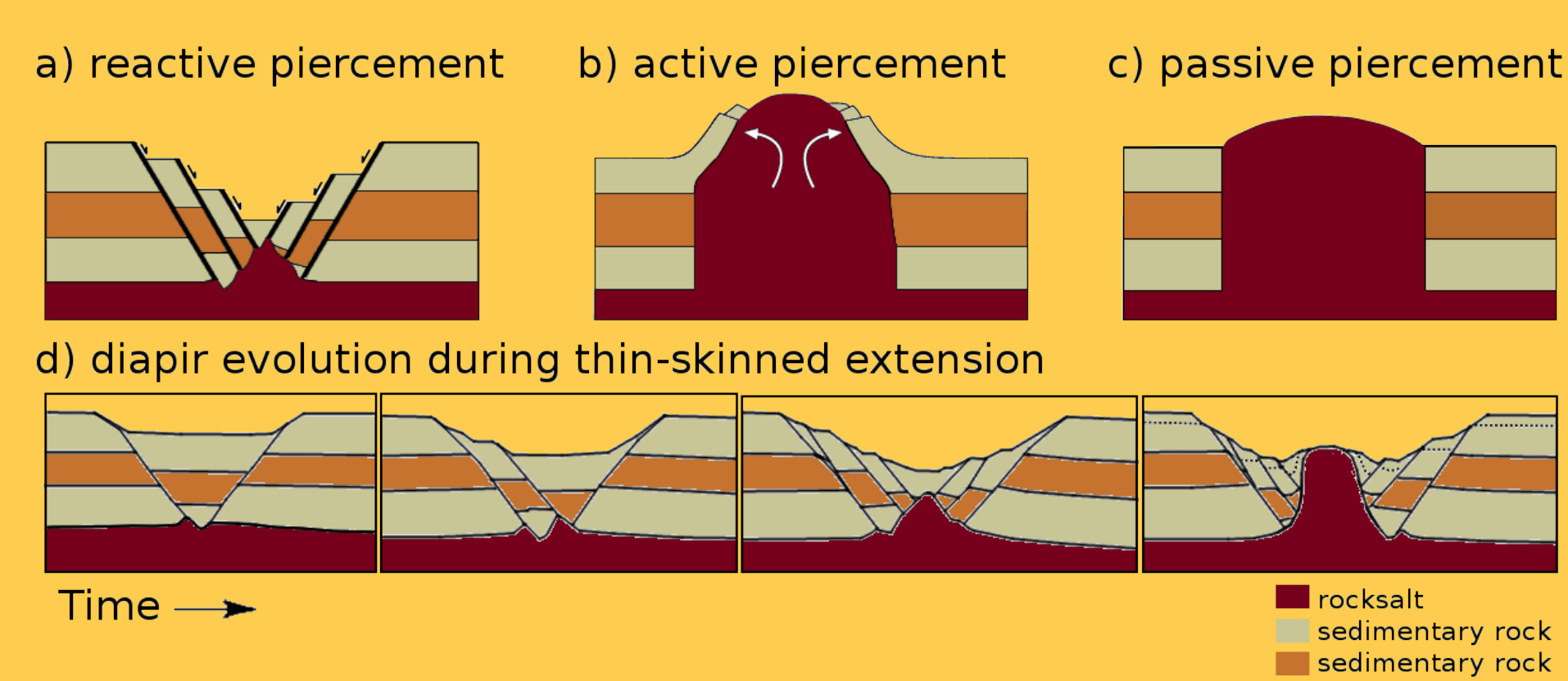
We assume that materials in the Earth's crust behave like incompressible visco-plastic fluids. Their flow is then best described by the Stokes equations:

$$\nabla(\mu(\nabla v + \nabla v^T)) - \nabla p + \rho g = 0 \quad (1)$$

$$\nabla \cdot v = 0 \quad (2)$$

Equation (2) is the mass conservation equation for incompressible fluids and equation (1) is the momentum conservation equation. This set of equations is solved by means of a Finite Element code similar to the one presented in [1].

3 - STRUCTURAL GEOLOGY



Different modes of diapirism (a,b,c), after [2] and the evolution of a diapir as results from analogue experiment (d), after [3].

7 - OUTLOOK

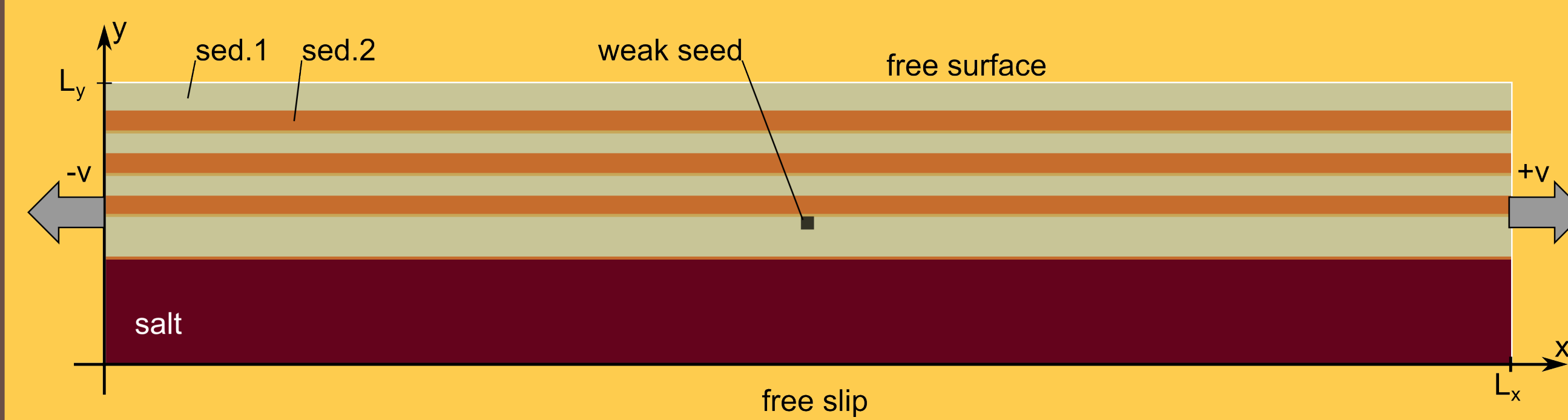
Points considered for future research:

- Nonlinear viscous rheology for salt
- Expand to 3D
- Inclined basement (often the drive behind thin-skinned extension)

REFERENCES

- [1] C. Thieulot, 2011. FANTOM: Two- and three-dimensional numerical modelling of creeping flows for the solution of geological problems In *Phys. Earth Planet. Inter.*
- [2] M. R. Hudec, M. P. A. Jackson, 2007. Terra infirma: Understanding salt tectonics In *Earth-Sci. Rev.*
- [3] B. C. Vendeville, M. P. A. Jackson, 1992. The rise of diapirs during thin-skinned extension In *Mar Pet Geol.*

4 - SETUP



$L_x = 20\text{km}$, $L_y = 4\text{km}$.
 Salt: linear viscous rheology, density $\rho_0 = 2100\text{ kg}\cdot\text{m}^3$, viscosity $\mu = 1 \times 10^{18}\text{ Pa}\cdot\text{s}$.

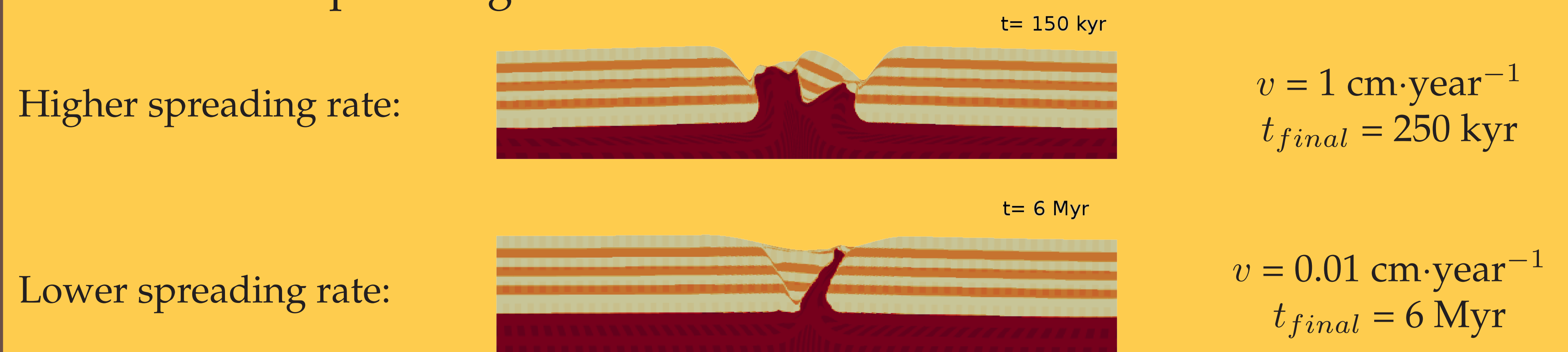
Sedimentary rocks: nonlinear plastic rheology, $\rho_0 = 2500\text{ kg}\cdot\text{m}^3$ (type 1) or $2400\text{ kg}\cdot\text{m}^3$ (type 2), $\mu_0 \approx 1 \times 10^{24}\text{ Pa}\cdot\text{s}$, cohesion $c = 55\text{ MPa}$ (1) or 62 MPa (2), angle of friction $\phi = 32^\circ$ (1) or 30° (2).

5 - RESULTS

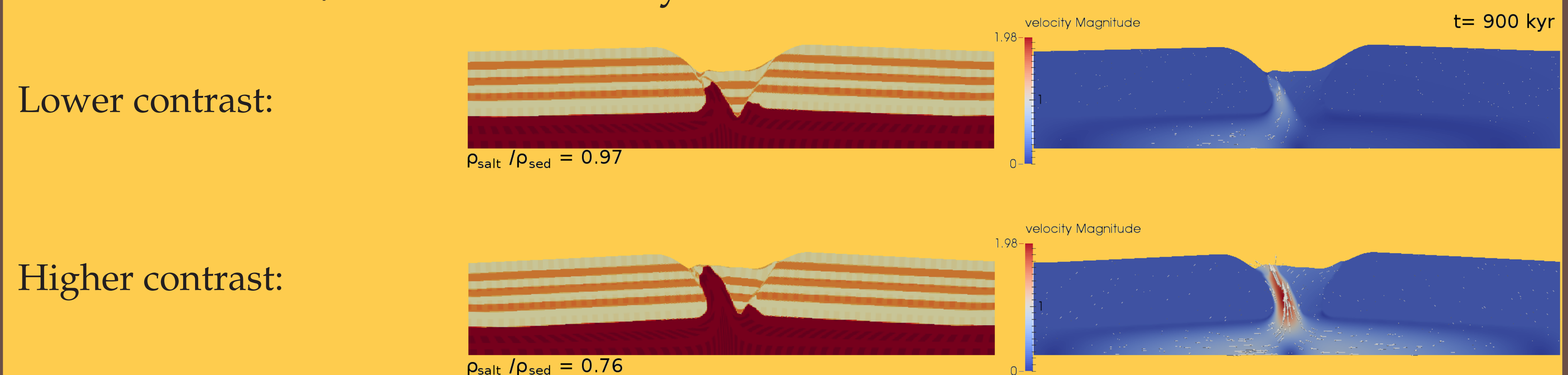
Evolution of reference model ($v = 0.1\text{ cm}\cdot\text{year}^{-1}$, $\rho_{\text{salt}}/\rho_{\text{sed}} = 0.86$, $t_{\text{final}} = 1\text{ Myr}$):



Differences in spreading rate:



Differences salt/sediment density contrast:



Differences in local surface processes vigour:



Perturbations on salt-sediment interface:



6 - CONCLUSIONS

Increased extension rates result in:

- Shorter surfacing time
- Relatively longer reactional phase
- Overall thicker diapir

Increased surface processes vigour results in:

- Longer surfacing time
- Decreased growth rate during active piercement

Increased density contrast results in:

- Shorter surfacing time
- Increased growth rate during active piercement

- Enhanced cone-shape

Enhanced salt-sediment interface perturbations results in:

- Enhanced asymmetry; half-grabens

Overall there is a good correlation with other structural-geological, analogue and numerical studies, as well as natural phenomena.