



Subduction initiation in the Neotethys: from field evidences to numerical modelling

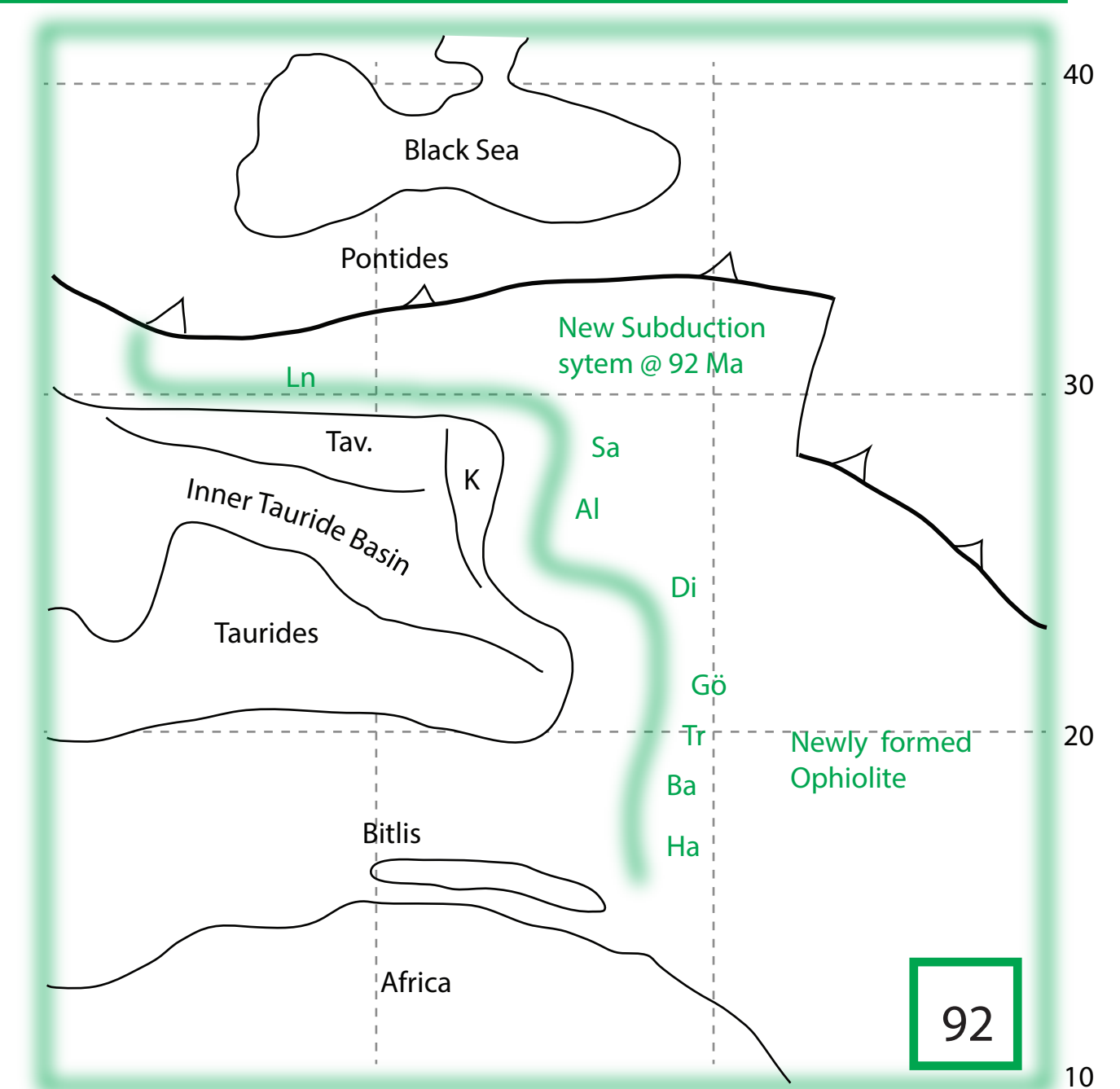
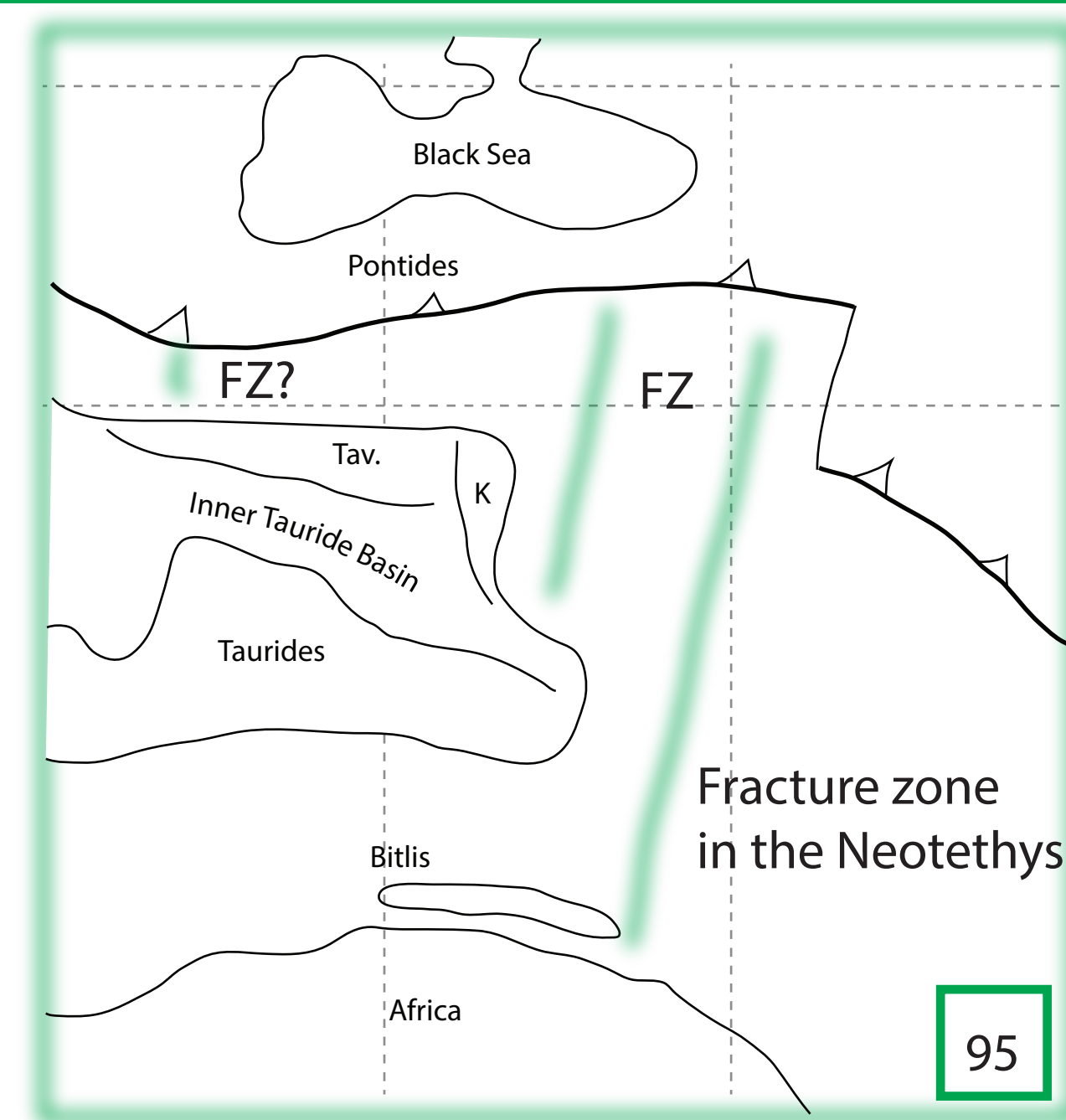
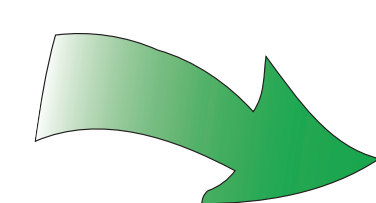
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{Introduction}

Subduction initiation is an unresolved question of plate tectonics. Few models have been proposed (either conceptual of physical; e.g. Stern 2004; Maffione et al., 2015; Leng & Gurnis 2015).

My approach

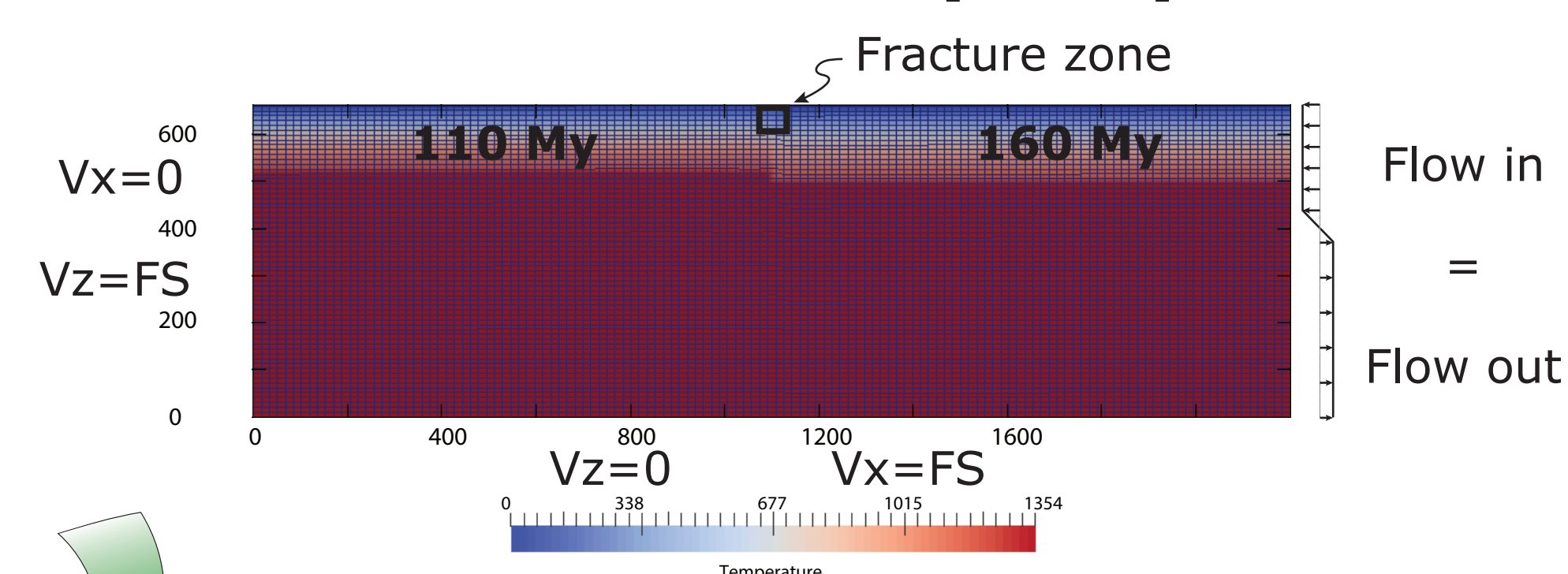
Here I chose to test whether it is possible to initiate subduction at fracture zones. Why? Because we have evidences that the ridge was gone since ca. 110 Ma in the Neotethys (in Turkey) where I did most of my previous work.



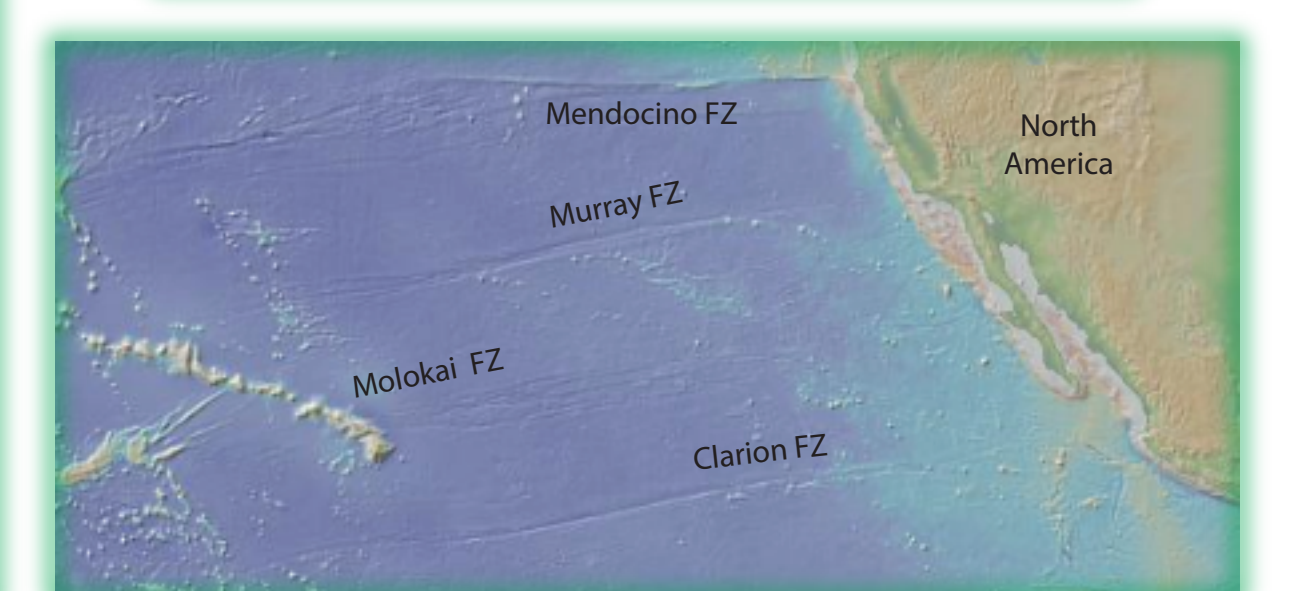
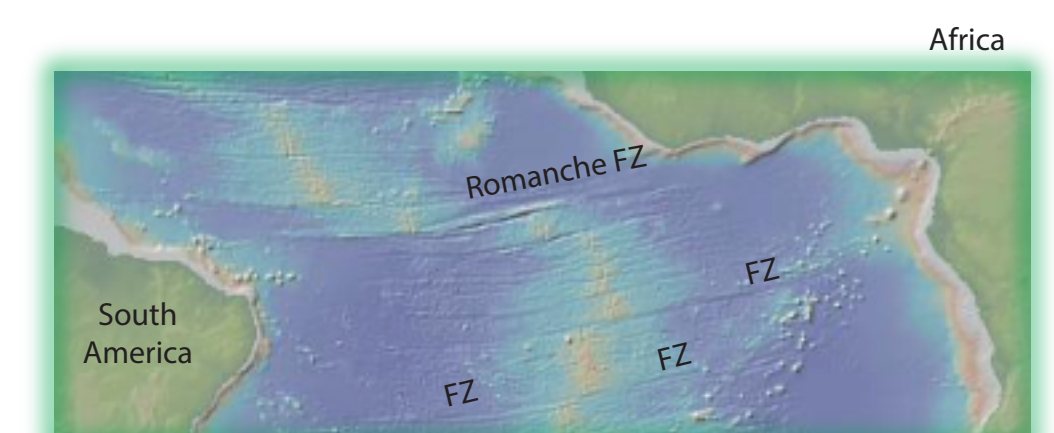
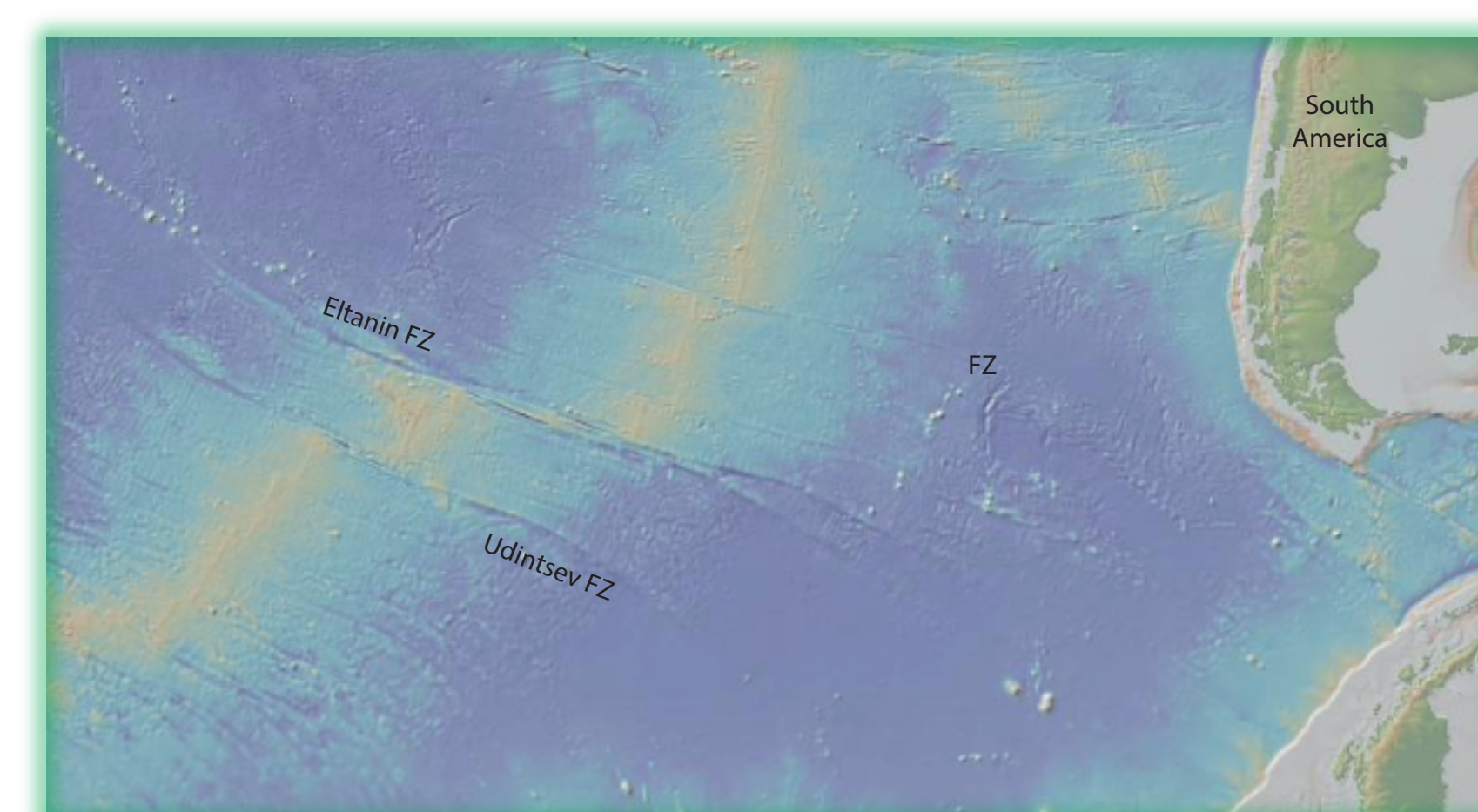
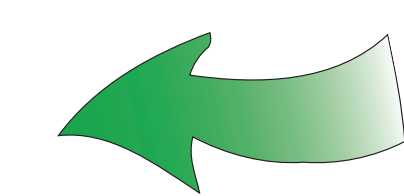
{ELEFANT} (Thieulot, 2014)

Finite element thermomechanically coupled code
Temperature profile of a ca. 110 vs. 160 My old oceanic lithosphere. Computed as a semi-infinite half-space.
Material: olivine with a non Newtonian and temperature dependant viscosity following Karato & Wu (1993)

$$\mu_{\text{eff}} = f \frac{1}{2} A^{\frac{1}{n}} \dot{\epsilon}_{\text{II}}^{\left(\frac{1}{n}-1\right)} \exp \left[\frac{Q + PV}{nRT} \right]$$



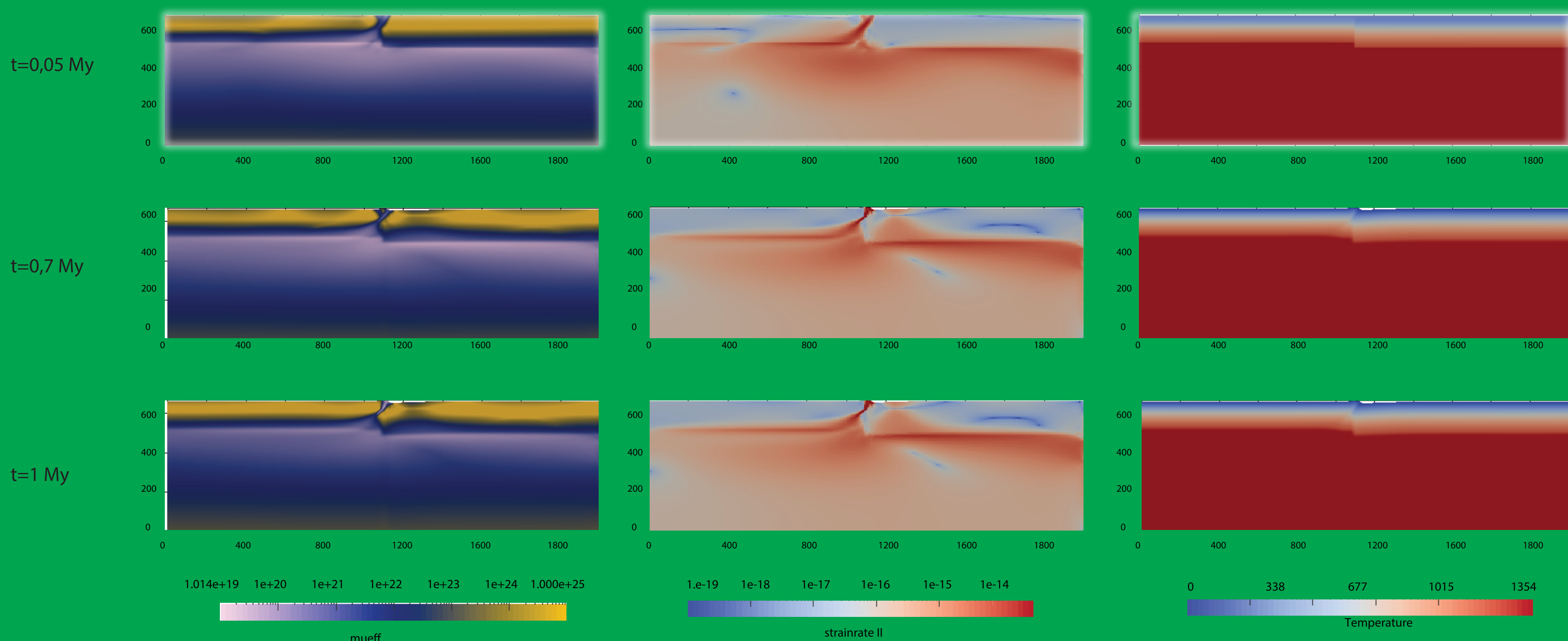
Present day configuration



Fracture zone are common objects. Average width is about 20 km and can go up to 100 km when there are close fracture zones (Romanche fracture zone for example).
Difference in age typically varies from 10 to 50 My.

{Preliminary results}

Subduction initiation at transform zones is possible!!!



To be tested

Possibility to form metamorphic sole
Influence of shear heating

Half out flow on the other side of the domain
Mesh refinement
Increased resolution

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