# Role of thermo-mechanical processes in seismicity: overview of STM modeling

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### Bridging time scales from tectonics to dynamic earthquake rupture

**1C** 1A Pro-foreland basin **1B** Retro-foreland basin Orogenic wedge 10 0 rate [m/s] Depth (km) ୍ଲି 20 back-thrusting 60 Depth Lower continental plate Upper continental plate 40  $10^{-3}$ 90 Slip 60 120  $10^{-8}$ 2150 1750 2050 2250 1650 1850 1950 80 50 60 70 80 90 100 Distance (km) Time [yr]



Utrecht University

110

#### **Tectonics**

- Thermo-mechanical evolution of convergent margins
- For millions years and over 100-1000's km's
- Static friction
- $-\Delta t = 1000$  years

- **Earthquake-like events on various faults:** - Rate-dependent friction
- Inertia added
- $-\Delta t = 1-5$  years
- **Rate-and-state dependent friction** - invariant reformulation w.  $\tau_{II} = \sigma_{yield}$ - From periodic earthquake cycles to slow and aseismic slip  $-\Delta t = milliseconds - years$

#### Dynamic earthquake rupture - Seismic wave propagation - Fault evolution $-\Delta t = milliseconds$

All done using same code based on

- 2D finite-differences on fully staggered grid w. marker-in-cell technique

A summary of what **seimo-thermo-mechanicanical models (STM)** can do:

- to solve conservation of mass, momentum and heat
- using a visco-elasto-plastic rheology and laboratory-based parameters for different lithologies
- with some hydro- (e.g., slab dehydration and fluid flow) and chemical components (e.g., serpentinization, melting)
- resolve interseismic, coseismic and postseismic phase
- resolve nucleation size and cohesive zone
- accurate and convergent solution on predefined fault
- grid size convergence through W=Wm\*log(1+K(V0/Vp))

See: Gerya & Yuen, PEPI, 2007; van Dinther et al., JGR, 2013a,b; GRL 2014; Dal Zilio et al., EPSL, 2018; Herrendoerfer et al., JGR, 2018; Preuss et al., EarthArXiV, 2019; in prep.

## How thermal coupling increases Mmax and b-value

>> Convergence rate not only increases seismic rate, but also maximum magnitude and proportion of large events (b)

(1) 10 million years of thermo-mechanical evolution of continent-continent collision leads to setup in Fig. 1A

(2) Spontaneous ruptures of various sizes and types are generated (see also Fig. 1B)



#### (3) Combined lead to Gutenberg-Richter law





# How thermo-mechanical coupling impacts seismicity

(1) How **sediment thickness** influences subduction dynamics and subduction zone seismicity





(3) Why modeling long-term dynamics is important

## How deep rheology and temperature affects seismicity

(1) Combine geological and geophysical data to simulate Northern Appenines (Italy)



### Lithosphere mechanics affects surface displacements

(1) STM predicts secondary zone of uplift

(4) Mechanism 2: Mass conservation following slab penetration

0.25

 $\mu_{s}$ 

0.50

long-term dynamics,

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no sedimenus (adapting friction) no long-term dynamics, no thermal coupling



### Summary

These examples illustrate how thermo-mechanical coupling, tectonics and rheology affect seismicity. More work and potential is being demonstrated by STM group at ETHZ/UU. Coupling with **fluids** is illustrated in the next poster by Petrini et al. Coupling with **chemisty** is illustrated in the presentation of Gerya et al. Meanwhile also developments in 3D and HTM coupling are nearing completion >>



<< 3D Long-term deformation

3D Dynamic earthquake rupture >>

Pranger et al., in prep.

