





Sensitivity Kernels for Receiver Functions -Prelimenary results from the Groningen Gas Field J.H.E. de Jong^{*}, H. Paulssen & J.A. Trampert

Receiver Functions Adjoint Sources

To assess the ability of receiver functions (Vinnik 1977, Langston 1979) to determine the depth of shallow layer boundaries in a regional, and possible time variable setting such as the Groningen gas-field, we have derived the adjoint sources (Tarantola 1984, Tromp et al. 2005) for a receiver function waveform misfit:



$$\chi = \frac{1}{2} \left[RF^{syn}(\mathbf{x}_r, m, t) - RF^{obs}(\mathbf{x}_r, t) \right]^2$$

We found the adjoint sources to be:

$$f_{R}^{\dagger}(\mathbf{x}, T-t) = F^{-1} \left[\frac{\Delta RF(\mathbf{x}_{r}, \omega)}{u_{V}^{*}(\mathbf{x}_{r}, \omega)} \right] \delta(\mathbf{x} - \mathbf{x}_{r})$$
$$f_{V}^{\dagger}(\mathbf{x}_{r}, T-t) = F^{-1} \left[\frac{\Delta RF(\mathbf{x}_{r}, \omega)u_{R}^{*}(\mathbf{x}_{r}, \omega)}{u_{V}^{2^{*}}(\mathbf{x}_{r}, \omega)} \right] \delta(\mathbf{x} - \mathbf{x}_{r})$$

We calculate the sensitivity kernels using synthetic data (figure 1) from 2D models of the Groningen Gas field (figure 3 & 4). This data is based on real test shots (figure 2) measured at a borehole near Stedum. In figure 1, we show the synthetic traces, the receiver functions, their difference, and the adjoint sources for the lowest reservoir geophone. In figure 2, we show the synthetic and real receiver functions for all 10 geophones in the reservoir.



Sensitivity Kernels

In figure 3 & 4, we show the sensitivity kernels and the 2Dmodel. The most important observations are:

- 1. Significant P-wave speed sensitivity (α-kernel) before conversion
- 2. Strong sensitivity to S-wave speed after conversion.
- 3. Sensitivity to reflections or conversions in the impedance kernel (ρ ').

Objective: Application in Groningen Context

Here we show our first, very preliminary, results. In the continuation of this work, we hope to investigate what information borehole receiver functions contain. Borehole receiver functions allow us to partly bypass both the unconsolidated upper layers, with all their challenges, as well as the source mechanism. Our results indicate that sensitivity kernels can help identify where conversions observed in receiver functions originate. We hope to see whether these kernels might be used to detect temporal variations within the reservoir or within the overburden. Using receiver functions, we can focus in particular on the depth of layer boundaries.





References:

(yellow), for the first arrival (see fig. 1).

Langston, C. A. Structure under Mount Rainier, Washington, inferred from teleseismic body waves. J. geophys. Res. Solid Earth, 120:537–543, 1979.
Tarantola, A. Inversion of seismic reflection data in the acoustic approximation. Geophysics, 49(8):1259–1266, 1984a.
Tromp, J., C. Tape, and Q. Liu. Seismic tomography, adjoint methods, time reversal and banana-doughnut kernels. Geophys. J. Int., 160:195–216, 2005.

L. Vinnik. Detection of waves converted from P to SV in the mantle. Phys. Earth Planet. Inter., 15:39–45, 1977.

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